Despite the rapid progress in angiosperm palaeobotany during recent decades, the ecology and life habits of early angiosperms are still poorly documented. There are several hypotheses concerning early angiosperm ecology and environments.

According to the traditional view, early angiosperms were similar to primitive woody magnoliids with evergreen leaves and large bisexual flowers and they inhabited wet forest understory (Takhtajan 1969, Thorne 1976).

A related view, based on new phylogenetic reconstructions and data on the morphology and ecophysiology of the most primitive extant plants, reconstructed early angiosperms as woody plants, similar to *Amborella*, which grew in wet, disturbed forest understory habitats or shady streamside settings (Field et al. 2004).

Others authors suggested that the early angiosperms were weedy xeric shrubs or riparian weeds that lived in open, disturbed habitats in semiarid areas or in disturbed streamside habitats in mesic environments (Stebbins 1965, Hickey and Doyle 1977). These hypotheses were based on arguments that variable conditions might have favored the evolution of the reproductive and vegetative features of angiosperms. This is supported by the findings of the first angiosperms in what were semiarid tropical and subtropical environments.

The palaeoherb hypothesis was based on the results of early phylogenetic analyses which placed angiosperms among Nymphaeales, Piperaceae, Saururaceae, Aristolochiaceae, and Chloranthaceae (Taylor and Hickey 1996). It inferred that the earliest angiosperms were ruderal, fast-growing herbs or shrubs, growing on sunny, unstable streamsides.

Sun and co-authors suggested an aquatic origin for angiosperms based of the life habits of *Archaefructus* and *Sinocarpus* (Sun et al. 2002, 2008). Freshwater plants are also in the angiosperm assemblage from the Chengzihe Formation (eastern Heilongjiang, China). The angiosperm remains are accompanied by the fern *Onychiopsis psilotoides* which is represented by almost entire young plants. The plant fossils were buried during a single flooding event and remained very close to their original location. They formed a pioneer open herbaceous community, consisting of ferns and angiosperms with a predominance of the latter and adapted to colonize fresh sediments in periodically flooded areas.

**Key words:** early angiosperms, palaeoecology, Early Cretaceous, Albian, Far East, Russia

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Unfortunately, most parts of these hypotheses are based on habits, morphology, and ecology of modern species and usually cannot be verified with fossil data.

Here we report on new finds from the Albian deposits of Primorye, Far East of Russia, which allows reconstruction of the ancient angiosperm dominated herbaceous communities and their role in the Early Cretaceous vegetation. The fossil plants were buried in almost autochthonous conditions during a single flooding event and remained very close to their original location. The locality of Bolshoy Kuvshin yielded numerous fern and herbaceous angiosperm remains with the latter prevailing. The quality of the material is exceptional and almost complete plants were often found.

This locality differs considerably from many other early angiosperm localities in Primorye and worldwide which usually contain an allochthonous combination of different Early Cretaceous plants with an insignificant admixture of small shrubby or herbaceous angiosperm. It yielded herbaceous fossils that are usually not preserved because of their delicate nature. The most autochthonous or paraautochthonous localities with large numbers of herbaceous angiosperms usually reflect an aquatic environment and occur significantly later in the geological record (Wang and Dilcher 2006, Cúneo et al. 2014). Autochthonous localities with a predominance of herbaceous terrestrial angiosperms are not known in the Lower Cretaceous and, as far as we know, neither in younger strata.

The fossils from Bolshoy Kuvshin have the potential to provide new data on the habits and environment of early angiosperms, important for understanding their initial radiation and also for investigation of vegetation structure and plant interactions in the Early Cretaceous.

In this paper we present a preliminary general description of the floristic composition and taphonomy of the locality Bolshoy Kuvshin and discuss in detail its stratigraphical position and age.

Material and methods

The locality Bolshoy Kuvshin is situated on the coast of the Ussuri Bay on the Bolshoy Kuvshin Cape in the Bolshoy Kamen Inlet near the town of Bolshoy Kamen in Primorye region, Far East of Russia (Text-figs 1, 2).

It was discovered by the geologist A. Oleynikov and then studied by palaeobotanists V. Krassilov and E. Volynets (Krassilov and Volynets 2008), who described two new species of tiny herbaceous angiosperms from there: *Achaenocarpites capitellatus* Krassilov et Volynets and *Ternaricarpites floribundus* Krassilov et Volynets.

Fossils come from an outcrop of black shales and greenish gray tuffaceous sandstones exposed along the railroad in the base part of the Bolshoy Kuvshin Cape (Text-fig. 3a). A detailed description is given in the stratigraphy section. The sequence commences and ends with thick layers of conglomerates. Specimens were found in a layer of fine-grained sandstone about 15 cm thick which is located 14 m above the lower conglomerate layer (Text-fig. 3b). The coordinates of the site are 43º07.853’N and 132º20.089’E.

The fossiliferous sandstone is massive, without pronounced bedding planes. The grain size of the sandstone changes significantly, but gradually, throughout the layer. This layer is underlain and overlain by finely laminated stratified sandstones. The plant remains were buried in the bedding planes, as well as folded or imbedded obliquely in the sediment, crossing the bedding planes. It is likely that this fossil-bearing layer was formed during one catastrophic flooding event.
A new excavation at this locality was made in summer 2017 by L. Golovneva, P. Alekseev, E. Bugdaeva and E. V olynets. The fossiliferous layer was uncovered in a square of about 3 m². About two hundred specimens were collected. All specimens are deposited in the palaeobotanical collection of the Federal Scientific Center of the East Asia Terrestrial Biodiversity FEB RAS in Vladivostok (former Institute of Biology and Soil Science), and are given the prefix IBSS 320-.

The plant fossils are preserved as brownish or dark gray impressions, yielding no structurally preserved material. Sometimes it is possible to recognize almost whole plants, with rhizomes, branching stems, leaves and fruits in organic connection. Although the preservation does not allow any study of anatomical details, these fossils provide important information on the overall habit and ecology of early angiosperms.

The material was studied under a Zeiss Stemi 2000-C binocular microscope, and photographed using a digital Nikon Coolpix P7700 camera at low-angle illumination or in water.

In this paper we assign the exposure on the Bolshoy Kuvshin Cape to the upper part of the Frentsevka Formation of the Partizansk coal basin. But the stratigraphical subdivision of the Lower Cretaceous deposits of this area is debatable. Other authors considered this sequence as the lower part of the Kangauz Formation (Markevich et al. 2000, Volynets 2005).

Stratigraphy

Non-marine Cretaceous sediments with subordinate marine or brackish deposits are widely distributed in Primorye. There are two big coal basins in the southern part of this region: the Partizansk (previously Suchan) coal basin and the Razdolnaya (previously Suifun) coal basin (Text-fig. 1). The Partizansk basin is situated along the Partizanskaya (previously Suchan) River valley, extending from the eastern coast of the Ussuri Bay to the upper reaches of the Ussuri River at a distance of about 120 km. The main industrial deposits are located in the eastern part of the basin, near Partizansk city. Here the Lower Cretaceous deposits have been studied most intensively.

The stratigraphy of the Lower Cretaceous strata in the Partizansk basin was studied by Kryshtofovich (Kryshtofovich 1929, Kryshtofovich and Prynada 1932), Perepechina (1960), Sharudo (1960), Vereshchagin (1977), Shtempel’ (1960), Krassilov (1967), Konovalov (Markevich 1980).

### Table 1. The stratigraphy of the Lower Cretaceous deposits in the Partizansk coal basin.

<table>
<thead>
<tr>
<th>Stage</th>
<th>Groups</th>
<th>Western part (eastern coast of the Ussuri Bay)</th>
<th>Eastern part (near Partizansk city)</th>
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<tbody>
<tr>
<td></td>
<td></td>
<td>Romanovka Fm</td>
<td>Kangauz Fm</td>
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<tr>
<td>Cenomanian</td>
<td>lower</td>
<td>Korkino Group</td>
<td>Unit of black siltstones</td>
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<tr>
<td></td>
<td>upper</td>
<td></td>
<td>Unit of heterogranaul sandstones</td>
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<td>middle</td>
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<tr>
<td></td>
<td>lower</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Aptian</td>
<td>upper</td>
<td>Suchan Group</td>
<td>Coal-bearing unit</td>
</tr>
<tr>
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<td>lower</td>
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<tr>
<td>Hauterivian</td>
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<tr>
<td>Valanginian</td>
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</table>
The coal-bearing deposits are together known as the Suchan Group, 1,300–1,700 m thick. It lies with erosion and angular unconformity on the Proterozoic gabbroic rocks or on the Klyuchi Formation (Tab. 1). This formation records shallow-marine to nonmarine environments with buchiid bivalves and plants of the Valanginian age (Markevich et al. 2000). The overlying variegated or red-colored volcaniclastic deposits make up the Korkino Group (up to 1,500 m), which is divided into the Kangauz (mostly sandstones) and the Romanovka (mostly siltstones) formations. This sequence records the initial phase of the Sikhote-Alin orogenic events.

The coal-bearing deposits of the Suchan Group yielded numerous fossil plants which were studied by Krassilov (1967). There are several alternative versions of the stratigraphic subdivision of the coal-bearing strata of the Partizansk coal basin. Here we use the scheme proposed by Likht (1961, 1994) and Krassilov (1967). According to these authors the Suchan Group consists of the Starosuchan (Neocomian), Severosuchan (Aptian) and Frentsevka (Albian) formations.

The upper boundary of the Severosuchan Formation is marked by the last thick coal layer known as the Velikan (Giant). The oldest occurrences of angiosperm come from the bottom of this layer where several *Araliaphyllum luciferum* (KRYSH.) GOLOVN. leaves were found (Kryshtofovich 1929, Golovneva 2018).

The Frentsevka Formation includes marine *Trigonia* beds, nonmarine black shales and a measure of greenish gray plant-bearing sandstones and siltstones (Text-fig. 4). *Trigonia* beds usually begin from a bedset (about 30
m) of fine-grained grey or greenish-grey heterogranular sandstones with the remains of trigonidiids (Quadratrotigonia (Transitrigonia) fusinensis Mirol., Pterotrigonia hokkaidoana (Yeh.), P. pocilliformis (Yok.), Ussuritrigonia subpyriformis Konov.), following by beach deposits with fragments of Isognomon sp., Lima sp., Ostrea sp., and Callista pseudoplena Yabe et Nagao, representing deposits of a short-term marine ingresson. The age of this assemblage was estimated as middle Albian (Markevich et al. 2000). It is overlain by fine-grained sandstones and siltstones with an admixture of conglomerates and rare thin coal layers. The thickness of this bedset is 70–80 m. These sediments include fossil plants Onychiopsis psilotoides (Stokes et Webb) Ward, Coniopteris burejensis (Zalessky) Seward, Cladophlebis frigida (Heer) Seward, C. novopokrovskii Prynada, Polyodontites polyisorus Pryn., Elatides asiatica (Yokoyama) Krassilov and Athrotaxopsis expansa Fontaine (Krassilov 1967). The overall thickness varies between 85 to 250 m. The Albian age of this measure was confirmed by the occurrence of Inoceramus concentricus Park in the south part of the Partizansk coal basin from the Vladimirskaya Unit (Oleynikov et al. 1998) which are considered as deposits from the same transgression (Markevich et al. 2000).

The black shales measure is 100–150 m thick and consists of interbedded siltstones and mudstones with beds of fine-grained sandstones, representing mostly lacustrine deposits. These sediments yields freshwater molluscs Limnocyrena andersonsoni (Grab.), Campeloma yihsiensis Grab., C. tani Grab., Viviparus cf. matumotoi Suz., ostracodes and conchostracans.

The overlying measure is 60–150 m thick and consists of intercalations of greenish-gray tuffaceous sandstones and siltstones with an admixture of coarse-grained sandstones and conglomerates. It yielded fossil plants Onychiopsis psilotoides, Coniopteris burejensis Podozamites sp., Elatides asiatica, E. ex gr. curvifolia (Dunker) Nath., Elatocladus sp. and others, including remains of angiosperms, represented by small Sapindopsis leaves. This sequence begins with conglomerates (about 1.5–45 m in thickness) which were interpreted as marking a regional erosion event (Markevich et al. 2000).

These deposits are conformably overlain by the volcaniclastic Korkino Group. This tuffaceous sequence contains very few fossil plants (Volynets 2005) and its age is inferred as late Albian – early Cenomanian mostly on the basis of its stratigraphical position.

Marine deposits with Trigonia are distributed only in the eastern part of the Partizansk coal basin. For this reason some authors did not consider the Frentsevka Formation as a separate stratigraphical unit. The Trigonia beds and nonmarine black shales were assigned to the upper subformation of the Severosuchan Formation, and the greenish gray plant-bearing sandstones were assigned to the Kangauz Formation of the Korkino Group (Perepechina et al. 1958, Sharudo 1960, Markevich et al. 2000).

However, Krassilov (1967) noted that the characteristics of the terrigenous sediments and composition of plant fossils in the Trigonia beds and measure of greenish gray plant-bearing sandstones and siltstones are similar, and interpreted the conglomerates as marking only local disconformities. In this article we follow Krassilov’s point of view on the size of the Frentsevka Formation. Its total thickness near Partizansk city is about 600 m.

The western part of the Partizansk coal basin, adjacent to the Ussuri Bay shore is called the Petrovka coal area or the Petrovka depression. The Cretaceous deposits here were also assigned to the Suchan and Korkino groups, but they are less thick and they do not contain industrial coal seams. The stratigraphical scheme of the Cretaceous deposits of this area was developed by Perepechina (1960). The deposits of the Suchan Group are divided into two formations: coal-bearing and overlaying coal-free formations. The latter was correlated with the Frentsevka Formation (Krassilov 1967) and was divided into two units: heterogranular sandstones and black siltstones (Tab. 1, Text-fig. 4). A significant section of these units is exposed along the southwestern shore of the Cape Palets, near the town of Bolshoy Kamen. The heterogranular sandstone unit is 300 m thick and is composed mainly of medium- and coarse-grained sandstones with numerous layers of fine-grained sandstones and siltstones, as well as layers of gravelites and conglomerates. There are occasional layers of coal-bearing sediments. This unit yielded Coniopteris sp., Yargolopteris rossica Pryn., Ruffordia goeppertii (Dunker) Seward, Lobifolia novopokrovskii (Pryn.) Rassk. et Lebedev, Elatides asiatica (Yokoyama) Krassilov, E. ex gr. curvifolia (Dunker) Nath. and Athrotaxopsis expansa Fontaine.

Text-fig. 5. a – unit of black siltstones from the Frentsevka Formation; b – small fish, scale bar 0.5 cm.
The unit composed of black siltstones is about 100 m thick. It begins and ends with conglomerates 10–30 m in thickness. Above and below them is interlayering of sandstones and siltstones. In the middle of this unit there are 30–40 m of black siltstones with numerous freshwater invertebrates and small fishes (Text-fig. 5).

Marine bivalves were not detected in this section in the Petrovka depression. The layer of black siltstones contains the same freshwater fauna as the measure of black shales near the city of Partizansk: *Sphaerium, Unio, Mycetopodus, Corbicula, Cyrena, Campeloma, Viviparum* (Vereshchagin 1977, Markевич et al. 2000).

The numerous plant fossils occur in all part of the Frentsevka Formation. In sandstones above lower conglomerates (Text-fig. 6) Krassilov discovered angiosperms *Araliaephylum (Sassafras) ussuriensis* (KRASSILOV) GOLOVN., *Artocarpidium* sp. and *Sapindopsis* sp. In addition to these also *Isoetes* sp., *Pelletieria ussuriensis* (PRYN.) KRASSILOV, *Onychiopsis psilotoides* (STOKES et WEBB) WARD, *Lobifolia novopokrovskii* (PRYN.) RASSK. et LEBEDEV, *Cycadites* sp., *Taeniopteris* sp., *Elatides asiatica* (YOKOYAMA) KRASSILOV, *E. ex gr. curvifolia* (DUNKER) NATL., *Atroraxops expansa* FONTAINE and *Brachyphyllum* ex gr. *obesum* HEER were found (Krassilov 1967).

The same section of black siltstones is exposed at the base of the Bolshoy Kuvshin Cape in the Bolshoy Kamen Inlet next to the railroad (Text-fig. 4). Krassilov (1967) collected *Coniopteris burejensis* (ZALESKY) SEWARD, *Onychiopsis psilotoides* (STOKES et WEBB) WARD, *Cladophlebis novopokrovskii* PRYN., *Atroraxops expansa* FONTAINE and *Elatocladus obtusifolia* OSHII in sandstones from this outcrop. Later in this outcrop a new site with herbaceous angiosperms was found (Krassilov and Volynets 2008). These sediments are conformably overlain by a measure of tuffaceous sandstones, assigned by Perepechina (1960) to the Kangaou Formation.

Detailed lithological and facies description of the Early Cretaceous sediments of the Partizansk coal basin is presented in the papers by Sharudo (1960) and Golozubov (1997). Deposits of the Frentsevka Formation reflect nonmarine sedimentation and are represented by alluvial-proluvial, lacustrine, paludal and coastal facies. The accumulation of sediments occurred within the lower part of an alluvial valley, transforming it into a marine bay or estuary. Near the Palets Cape and the Bolshoy Kuvshin Cape, alluvial-lacustrine deposits are intercalated with coarse-grained sandstones and conglomerates, representing deposits of streams flowing down from the upland, located above the position of Ussuri Bay and the Muravyov-Amursky Peninsula. The Korkino Group is represented mainly by alluvial facies.

**Plant assemblage**

During excavation of the fossiliferous layer in the Bolshoy Kuvshin locality in 2017 about two hundred plant fossils were found. Among them remains of angiosperms predominate (about 90 %). The majority of specimens are represented by fragments of branching stems with attached leaves or fruits that belonged to small herbaceous plants. Several almost complete plants were also found (Text-figs 7–9).


Among the angiosperms, remains of *Achaenocarpites capitellatus* occur most often. This species was described by Krassilov and Volynets (2008) based on fragmented material. The most complete specimen of *Achaenocarpites* is represented by a whole plant about 10 cm high with a straight thin branching stem, several whors of leaves and three terminal heads, consisting of numerous achenes (Text-fig. 7a). The leaves are stipulate, ternate, and pinnatisect. Krassilov compared this species with different representatives of Ranunculaceae.

The second most abundant species is *Asiatifolium elegans* (Text-figs 7f, 8b, c). Usually this species is represented by the upper parts of stems with several helically attached and closely spaced leaves. The leaves are entire-margined, very diverse both in shape and size. Usually they are oblong, lanceolate or obovate, with an obtuse to rounded apex and decurrent base. Venation is pinnate, brochidodromous; with 5–8 secondary veins. Fructifications, associated with these leaves, have not been detected. This species was first described from the Chengziche Formation, exposed near the city of Jixi in Northeastern China (Sun and Dilcher 2002). The systematic position of *Asiatifolium* is uncertain.

Remains of *Ternaricarpites floribundus* are also abundant. This species was described by Krassilov and Volynets (2008). The most complete specimen is represented by a slender branching axis with several fruits, consisting of two to five follicles, most commonly three (Text-fig. 9a, b). The leaves of this plant are unknown. Krassilov believed that the leaves belonged to the *Ternaricarpites* plant (Krassilov and Volynets 2008), but a more detailed study of the material showed that in fact the leaves were not attached to the stem. On the basis of morphological characters and comparisons with other fossils, it appears that *Ternaricarpites* may also be related to Ranunculaceae, such as *Achaenocarpites* (Krassilov and Volynets 2008). In addition it has some similarity with *Hyrcantha* (*Sinocarpus*) *decussata* (LENG et E.M.FRIS) DILCHER, G.SUN, Q.JI et H.Q.LI from the Barremian to Aptian Yixian Formation of northeastern China (Leng and
Text-fig. 7. Plant fossils from Primorye, Partizansk coal basin, Frentsevka Formation, Bolshoy Kuvshin locality, early – middle Albian.
a, c, e – *Achaenocarpites capitellatus* **Krassilov** et **Volynets**: a – spec. IBSS 320-132, c – spec. IBSS 320-132, e – spec. IBSS 320-120;
b – *Onychiopsis psilotoides* (Stokes et Webb) Ward, spec. – IBSS 320-165; d, g – branching infructescence with several follicular fruits:
d – spec. IBSS 320-145, g – IBSS 320-145; f – *Asiaefolium elegans* G.Sun, S.X.Guo et Shao L.Zheng, spec. IBSS 320-75. Scale bar 0.5 cm.
Friis 2003, 2006, Dilcher et al. 2007). The Chinese plant has similar slender branching stems with terminal fruits consisting of two – four carpels.

*Jixia pinnatipartita* has simple deeply pinnately lobed entire-margined leaves (Text-fig. 9c). The lobes are thin, sublinear, usually with additional small lobes and widened bases, decurrent up and down along the midvein. The leaf base is truncate or with triangular incision. Venation is pinnate, craspedodromous. Remains of this species are quite rare. As with *Asiatifolium*, this species was first described from the Chengziche Formation (Sun and Dilcher 2002). Fructifications, associated with this plant, and its systematics are not known.

In addition to the previous species, there are two additional finds that we considered as belonging to new species. One new species is represented by a swollen curved branching stem with an axillary bud and three rather large petiolate, suborbicular to ovate leaves up to 6 cm long, with entire or crenulate margin and pinnate, brochidodromous venation (Text-fig. 8a). The leaves are thin and often folded, which indicate an aquatic or semiaquatic habit. This plant closely resembles well-preserved nymphaeaceous *Pluricarpellatia peltata* B.MOH, BERNARDES-DE-OLIVEIRA et D.W.TAYLOR, described from the Aptian – lower Albian Crato Formation in Brazil (Mohr et al. 2008).

The other new species is represented by a branching infructescence with several follicular fruits, consisting of two-three carpels (Text-fig. 7d, g). This plant resembles *Ternaricarpites*, but its carpels are bigger, usually in pairs and basally fused.

The angiosperm remains are accompanied by relatively abundant fern remains (about 10% of specimens) and occasional Equisetum rhizomes and scale-leaved conifer shoots. The conifer component consists of fragmentary twigs (about 2–5 cm) of Athrotaxopsis expansa Fontaine and small Sequoia-like shoots. Ferns are represented mostly by Onychiopsis psilotoides (Stokes et Webb) Ward. The remains of other ferns are very fragmented. Among them Ruffordia goepertii (Dunker) Seward, Teilhardia tenella (Pryn.) Krassilov and Coniopteris sp. can be recognized. All these species are represented only by small pieces of fronds, no more than a few pinnules.

Leaves of Onychiopsis psilotoides, on the contrary, are usually almost complete. One specimen is represented by

whole plants, with leaves situated in several different planes in the sediment (Text-fig. 7b). Onychiopsis was widely distributed in the Early Cretaceous floras of Laurasia. This fern has leathery pinnules and xeromorphic sporangia, enclosed in an indehiscent envelope. Both features indicate water stress in the environment which corresponds to dwelling in open, sometimes brackish habitats (Friis and Pedersen 1990). Onychiopsis psilotoides is one of the most characteristic ferns in all strata of the Partizansk coal basin, which are mostly nonmarine. Perhaps, this species could also be a stream-side dweller.

Specimens of Onychiopsis from the Bolshoy Kuvshin locality differ from specimens of this fern from other sections in its smaller size. The length of the whole frond (without stalk) varies from 7 cm up to 30 cm, usually being about 10–12 cm. Leaves from other sites in the Partizansk basin usually reach 70–90 cm in length. This implies that in the Bolshoy Kuvshin predominantly young plants were buried.

**Discussion**

The age of the Frentsevka Formation was based mainly on marine molluscs from Trigonia beds. Konovalov estimated the age to be middle Albian (Markevich et al. 2000). But in reality the bivalve assemblage includes many endemic species. Other species with a wider distribution (Pterotrignia hokkaidoana, P. pocilliformis, Inoceramus concentricus) have a rather wide stratigraphical range during the Albian (Tashiro and Matsuda 1983, Matsukawa et al. 1997).

Freshwater molluscs, ostracodes and conchostracans from black shales are in general insufficiently studied. At present they suggest an age range from the Aptian to the Albian (Markevich et al. 2000).

But these deposits contain a rich plant assemblage including ferns, gingoaoleans, cycadophytes, czekanowskialeans, conifers and angiosperms. Krassilov (1967) noted that plant assemblages from all formations of the Partizansk coal basin have a similar composition and differ mainly in the quantitative ratio between species. The most abundant species in the Frentsevka Formation are Onychiopsis psilotoides (Stokes et Webb) Ward, Cladophlebis frigida (Heer) Seward, Elatides asiatica (Yokoyama) Krassilov, Athrotaxopsis expansa Fontaine, Coniopteris burejensis (Zalesky) Seward, Cladophlebis frigida (Heer) Seward, Lobifolia novapokrovskii (Pryn.) Rass. et Lебедев, Polygodites polysorus Pryn. Elatides asiatica (Yokoyama) Krassilov, Elatocladus ex gr. curvifolia (Dunker) Nath. and Athrotaxopsis expansa Fontaine. All these species have a wide distribution in the Early Cretaceous.

Volynets (2005, 2006) developed a very detailed stratigraphical sequence of the floristic assemblages for the Partizansk, Radzonlaya and Alchan coal basins. She distinguished two successive assemblages dated from middle to late Albian in the deposits of the Frentsevka Formation. Unfortunately these assemblages were not compared with any accurately dated assemblages in other regions. Therefore they are not of use for age determination of the Bolshoy Kuvshin locality.

Plants having the greatest biostratigraphic significance are early angiosperms. In the mid-Cretaceous angiosperms were already quite numerous and the majority of species were characterized by a narrow stratigraphical range. Unfortunately, many species were endemic which thus prevents any correlation with different basins. In addition to herbaceous plants from the Bolshoy Kuvshin locality, three species of angiosperms were recorded from the Frentsevka Formation: Araliaephylum (Sassafras) ussuriensis (Krassilov) Golovn., Artocarpidium sp., Sapindopsis cf. angusta (Heer) Seward et V.Conway.

In the extensively studied sequence of the Potomac Group in the eastern United States, Araliaephylum and Sapindopsis were recorded in the deposits referred to the lower part of the pollen Zone II-B (locality Bank near Brook). The age of this locality was determined as early to middle Albian (Doyle and Hickey 1976) or as middle to late Albian (Hochuli et al. 2006, Doyle and Upchurch 2014). In Zone I (Aptian to earliest Albian) only simple unlobed leaves with predominantly pinnate venation were recovered. From Zone II-C (latest Albian-Cenomanian) large platanoid leaves became abundant in channel and levee facies (Doyle and Upchurch 2014). In accordance with these data, the age of the Frentsevka Formation should be estimated as middle Albian, possibly including the beginning of late Albian.

However, the Bolshoy Kuvshin assemblage includes two species, in common with the flora of the Chengzihe Formation, including in Jixi Group: Jixia pinnatifartita and Asiatifolium elegans. The angiosperms from the Chengzihe Formation were initially considered to be Hauterivian – early Barremian in age (Sun and Dilcher 2002). More recently reinvestigation of Aucellina bivalves from the Jixi group indicated a younger age, from the Barremian to Albian (Gu et al. 1997). The Chengzihe Formation is dominated by nonmarine coal-bearing deposits that intercalated with several marine beds. It contains plant megafossils, ostracodes, fishes, reptiles, abundant bivalves and dinoflagellates (Sha et al. 2003). Nowadays the Chengzihe Formation is considered to be mainly Aptian but lower marine beds perhaps extends downward into the Barremian. The overlying Muling Formation is composed of nonmarine coal-bearing deposits intercalated rarely with brackish sediments. It is dominated by grey and greenish-grey fine-grained sandstones and siltstones interbedded with mudstones, tuffs and coal beds. The fossils recovered from the Muling Formation (brackish bivalves and dinoflagellate assemblage) suggest an Aptian – early Albian age (Sha et al. 2003).

Angiosperms were collected from the middle and upper parts of the Chengzihe Formation and are expected to be Aptian in age. Until the present time, early angiosperms were considered as a rapidly evolving group whose species have the narrow stratigraphical ranges. The difference in age of the same species from Primorye and Heilongjiang implies that Jixia pinnatifartita and Asiatifolium elegans had wider stratigraphical ranges, maybe from Aptian to early – middle (late?) Albian, or that the age determinations were not accurate enough and additional stratigraphical investigations are necessary. Perhaps, an early Albian age could be a possibility for both plant assemblages.

Overall, we accept an early – middle Albian age for early angiosperms from the Frentsevka Formation.
Deposits of the Frentsevka Formation near the Bolshoy Kuvshin Cape are represented by alluvial-lacustrine floodplain facies intercalated with coarse-grained sandstones and conglomerates, representing deposits of braided rivers flowing down from the upland where is now located the Ussuri Bay (Sharudo 1960).

The plant-bearing layer is about 15 cm thick and represented by fine-grained non-stratified sandstone. This layer is underlain and overlain by finely laminated sandstones and siltstones. Most likely, the latter were deposited on the floodplain as a result of the activity of small, meandering channels. The plant remains are often folded, rolled and cross bedding planes. It is likely that this fossil-bearing layer was formed during one single catastrophic flooding event. The locality contains numerous remains of diverse herbaceous angiosperms with an admixture of the fern *Onychiopsis psilotoides*. Many specimens are represented by almost complete plants or big fragments with attached leaves and fruits. Plants were small (10–30 cm high) and very delicate. Their complete preservation in rather coarse sediment indicates the absence of water transport. They were deposited very close to their original location. Thus, this locality can be considered as autochthonous, reflecting the vegetation which grew near to the burial place. We interpret this vegetation as an open herbaceous community, consisting of ferns and angiosperms with a predominance of the latter. This community occupied low flat plains between river channels. Conifers and other woody plants were perhaps excluded from this environment by periodic flood events, low drainage, and by the unstable groundwater level. In the plant-bearing layer conifers are represented mostly by very rare, small (2–5 cm) shoots of *Athrotaxopsis expansa*. These remains are considered as an allochthonous admixture in this locality, as well as the small rare fragments of the ferns *Ruffordia goeppertii*, *Teilhardia tenella* and *Coniopteris* sp.

In several meters below the plant-bearing layer there is leaf mat in the sandstone, consisting exclusively of *Athrotaxopsis expansa* twigs and cones. Possibly this plant can have formed thickets on levees and other elevated parts of a floodplain.

The fern *Onychiopsis psilotoides* is represented by almost complete, but small young plants. This implies that the fern-angiosperm herbaceous community was a pioneer, reflecting the early succession stage and adapted to colonize fresh sediments in periodically flooded areas. Powerful floods, bearing abundant sediments or migration of channels can periodically destroy pioneer communities. Krassilov suggested that herbaceous angiosperms could also colonize fresh ash fallouts (Krassilov and Volyntets 2008). Reconstruction of such an angiosperm herbaceous community is shown in Text-fig. 10.

The locality Bolshoy Kamen, containing abundant remains of several herbaceous angiosperm species, is significantly different from other early – middle Albian angiosperm sites, where the angiosperm can be diverse and well-preserved, but their remains occur rarely and irregularly.
The preservation of complete herbaceous angiosperm, ferns and gnetaleans is typical for the Aptian – early Albian Crato Formation in Brazil (Mohr and Friis 2000, Mohr et al. 2015), and for the Aptian Yixian Formation in Northeast China (Sun and Dilcher 1997, Sun et al. 1998, Leng and Friis 2003, Dilcher et al. 2007, Sun et al. 2008, Yang et al. 2013).

In the Yixian Formation herbaceous plant fossils occur in finely laminated lacustrine tuffaceous siltstones together with diverse invertebrate lake fauna and fishes (Pan et al. 2013).

The Crato plattenkalk limestones were also formed by a large freshwater or brakish-water lake. It was suggested that periodic torrent rainfall could have washed down into the water almost complete plants which grew not far from the lake shore (Mohr et al. 2015). Angiosperms constituted a minor scattered element of the nearshore vegetation (Mohr and Friis 2000). It is interesting that many herbaceous remains from the Crato Formation are represented by young plants, documenting their early growth stages (Rydin et al. 2003). It is possible that these plants also could have inhabited periodically disturbed places along streams or lake edge.

An authochthonous plant assemblage from the riparian deposits was described from the Bajo Grande locality, coming from the lower Aptian Anfiteatro de Tico Formation, Argentina (Cladera et al. 2007). This assemblage includes the liverwort Ricciopsis gradensis Cladera et al., ferns Adiantopteris tripinnata Cladera et al., and undetermined Schizaceae, and the gnetophyte Ephedra verticillata Cladera et al. Taphonomical studies suggest that these plants grew and were buried near the levee. This locality documents an open fern herbaceous vegetation with gnetaleans, but unfortunately without angiosperms.

Our data support Hickey and Doyle’s (1977) interpretation of early angiosperms as riparian weeds. They also show that early angiosperm preferred open wet fern communities, inhabited periodically flooded river valleys and coastal plains. Such environments were favorable for appearance of different aquatic and semiaquatic life forms, which are so numerous among the early angiosperms.

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